

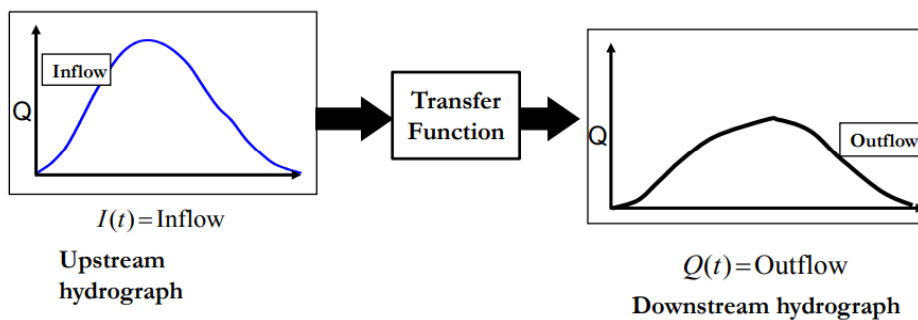
## Flow Routing Methods

Abhinav Shukla  
Scientist/Engineer 'SD'  
  
Disaster Management  
Support Group

### What is FLOW ROUTING?

- ✓ "Flow routing is a technique of determining the flow hydrograph at a section of a river by utilizing the data of flood flow at one or more upstream sections."

( Subramanya, 1984)



## Types of flow routing

### Lumped/hydrologic:

- Flow  $\rightarrow$  f(time)
- Continuity equation and Flow/Storage relationship

### Distributed/hydraulic:

- Flow  $\rightarrow$  f(space, time)
- Continuity and Momentum equations

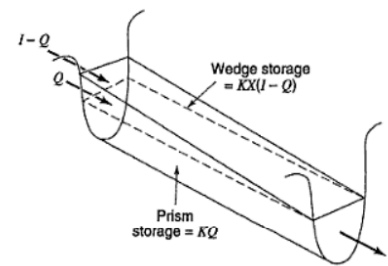
## Hydrologic flow routing

### Channel Routing

The total volume in storage for a channel reach having a flood wave can be considered as **prism storage + wedge storage**.

**Prism storage:** The volume that would exist if uniform flow occurred at the downstream depth i.e. the volume formed by an imaginary plane parallel to the channel bottom drawn at the outflow section water surface.

**Wedge storage:** It is the wedge like volume formed between the actual water surface profile and the top surface of the prism storage.



Prism and wedge storage in a channel reach (Mays, 2009)

### Hydrologic river routing (Muskingum Method)

Wedge storage in reach

$S_{Prism} = KQ$   
 $S_{Wedge} = KX(I - Q)$

$K$  = travel time of peak through the reach  
 $X$  = weight on inflow versus outflow ( $0 \leq X \leq 0.5$ )  
 $X = 0$  → Reservoir, storage depends on outflow, no wedge  
 $X = 0.0 - 0.3$  → Natural stream

$S = KQ + KX(I - Q)$   
 $S = K[XI + (1 - X)Q]$

Advancing Flood Wave  $I > Q$

Receding Flood Wave  $Q > I$

### Muskingum Method (Cont.)

$S = K[XI + (1 - X)Q]$   
 $S_{j+1} - S_j = K\{[XI_{j+1} + (1 - X)Q_{j+1}] - [XI_j + (1 - X)Q_j]\}$

Recall:

$$S_{j+1} - S_j = \frac{I_{j+1} + I_j}{2} \Delta t - \frac{Q_{j+1} + Q_j}{2} \Delta t$$

Combine:

$Q_{j+1} = C_1 I_{j+1} + C_2 I_j + C_3 Q_j$

$$C_1 = \frac{\Delta t - 2KX}{2K(1 - X) + \Delta t}$$

$$C_2 = \frac{\Delta t + 2KX}{2K(1 - X) + \Delta t}$$

$$C_3 = \frac{2K(1 - X) - \Delta t}{2K(1 - X) + \Delta t}$$

If  $I(t)$ ,  $K$  and  $X$  are known,  $Q(t)$  can be calculated using above equations

# Muskingum - Example

- Given:
  - Inflow hydrograph
  - $K = 2.3$  hr,  $X = 0.15$ ,  $\Delta t = 1$  hour, Initial  $Q = 85$  cfs
- Find:
  - Outflow hydrograph using Muskingum routing method

Period (hr)	Inflow (cfs)
1	93
2	137
3	208
4	320
5	442
6	546
7	630
8	678
9	691
10	675
11	634
12	571
13	477
14	390
15	329
16	247
17	184
18	134
19	108
20	90

$$C_1 = \frac{\Delta t - 2KX}{2K(1-X) + \Delta t} = \frac{1 - 2 * 2.3 * 0.15}{2 * 2.3(1 - 0.15) + 1} = 0.0631$$

$$C_2 = \frac{\Delta t + 2KX}{2K(1-X) + \Delta t} = \frac{1 + 2 * 2.3 * 0.15}{2 * 2.3(1 - 0.15) + 1} = 0.3442$$

$$C_3 = \frac{2K(1-X) - \Delta t}{2K(1-X) + \Delta t} = \frac{2 * 2.3 * (1 - 0.15) - 1}{2 * 2.3(1 - 0.15) + 1} = 0.5927$$

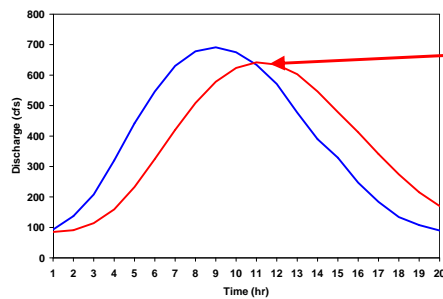
7

# Muskingum – Example (Cont.)

$$Q_{j+1} = C_1 I_{j+1} + C_2 I_j + C_3 Q_j$$

$C_1 = 0.0631$ ,  $C_2 = 0.3442$ ,  $C_3 = 0.5927$

Period (hr)	Inflow (cfs)	$C_1 I_{j+1}$	$C_2 I_j$	$C_3 Q_j$	Outflow (cfs)
1	93	0	0	0	85
2	137	9	32	50	91
3	208	13	47	54	114
4	320	20	72	68	159
5	442	28	110	95	233
6	546	34	152	138	324
7	630	40	188	192	420
8	678	43	217	249	509
9	691	44	233	301	578
10	675	43	238	343	623
11	634	40	232	369	642
12	571	36	218	380	635
13	477	30	197	376	603
14	390	25	164	357	546
15	329	21	134	324	479
16	247	16	113	254	413
17	184	12	85	245	341
18	134	8	63	202	274
19	108	7	46	162	215
20	90	6	37	128	170



8

**EXAMPLE 8.5** Route the following flood hydrograph through a river reach for which  $K = 12.0$  h and  $x = 0.20$ . At the start of the inflow flood, the outflow discharge is  $10 \text{ m}^3/\text{s}$ .

Time (h)	0	6	12	18	24	30	36	42	48	54
Inflow ( $\text{m}^3/\text{s}$ )	10	20	50	60	55	45	35	27	20	15

**SOLUTION:** Since  $K = 12$  h and  $2Kx = 2 \times 12 \times 0.2 = 4.8$  h,  $\Delta t$  should be such that  $12 \text{ h} > \Delta t > 4.8$  h. In the present case  $\Delta t = 6$  h is selected to suit the given inflow hydrograph ordinate interval.

Using Eqs. (8.16-a, b & c) the coefficients  $C_0$ ,  $C_1$  and  $C_2$  are calculated as

$$C_0 = \frac{-12 \times 0.20 + 0.5 \times 6}{12 - 12 \times 0.2 + 0.5 \times 6} = \frac{0.6}{12.6} = 0.048$$

$$C_1 = \frac{12 \times 0.2 + 0.5 \times 6}{12.6} = 0.429$$

$$C_2 = \frac{12 - 12 \times 0.2 - 0.5 \times 6}{12.6} = 0.523$$

For the first time interval, 0 to 6 h,

$$I_1 = 10.0 \qquad C_1 I_1 = 4.29$$

$$I_2 = 20.0 \qquad C_0 I_2 = 0.96$$

$$Q_1 = 10.0 \qquad C_2 Q_1 = 5.23$$

$$\text{From Eq. (8.16)} \quad Q_2 = C_0 I_2 + C_1 I_1 + C_2 Q_1 = 10.48 \text{ m}^3/\text{s}$$

For the next time step, 6 to 12 h,  $Q_1 = 10.48 \text{ m}^3/\text{s}$ . The procedure is repeated for the entire duration of the inflow hydrograph. The computations are done in a tabular form as shown in Table 8.4. By plotting the inflow and outflow hydrographs the attenuation and peak lag are found to be  $10 \text{ m}^3/\text{s}$  and 12 h respectively.

**Table 8.4** Muskingum Method of Routing—Example 8.5

$\Delta t = 6 \text{ h}$					
Time (h)	$I$ ( $\text{m}^3/\text{s}$ )	$0.048 I_2$	$0.429 I_1$	$0.523 Q_1$	$Q$ ( $\text{m}^3/\text{s}$ )
1	2	3	4	5	6
0	10				10.00
6	20	0.96	4.29	5.23	10.48
12	50	2.40	8.58	5.48	16.46
18	60	2.88	21.45	8.61	32.94
24	55	2.64	25.74	17.23	45.61
30	45	2.16	23.60	23.85	49.61
36	35	1.68	19.30	25.95	46.93
42	27	1.30	15.02	24.55	40.87
48	20	0.96	11.58	21.38	33.92
54	15	0.72	8.58	17.74	27.04

## Hydrologic flow routing

### Modified Pul's Method

$$I - Q = \frac{dS}{dt}$$

$$\bar{I} \Delta t - \bar{Q} \Delta t = \Delta S$$

$$\left( \frac{I_1 + I_2}{2} \right) \Delta t - \left( \frac{Q_1 + Q_2}{2} \right) \Delta t = S_2 - S_1$$

$$\left( \frac{I_1 + I_2}{2} \right) \Delta t + \left( S_1 - \frac{Q_1 \Delta t}{2} \right) = \left( S_2 + \frac{Q_2 \Delta t}{2} \right)$$

## Flow routing in channels

- Distributed Routing
- St. Venant equations
  - Continuity equation

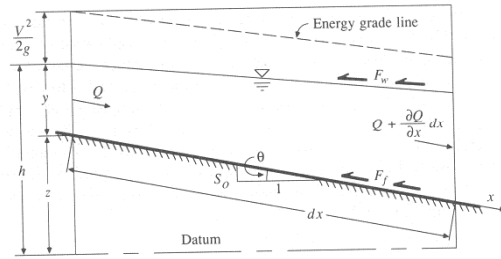
$$\frac{\partial Q}{\partial x} + \frac{\partial A}{\partial t} = 0$$

- Momentum Equation

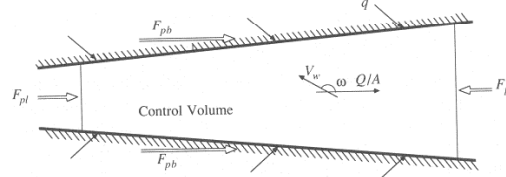
$$\frac{1}{A} \frac{\partial Q}{\partial t} + \frac{1}{A} \frac{\partial}{\partial x} \left( \frac{Q^2}{A} \right) + g \frac{\partial y}{\partial x} - g(S_o - S_f) = 0$$

What are all these terms, and where are they coming from?

# Continuity Equation



Elevation View



Plan View

$Q$  = inflow to the control volume

$q$  = lateral inflow

$\frac{\partial Q}{\partial x}$  Rate of change of flow with distance

$Q + \frac{\partial Q}{\partial x} dx$  Outflow from the C.V.

$\frac{\partial(\rho A dx)}{\partial t}$  Change in mass

Reynolds transport theorem

$$0 = \frac{d}{dt} \iiint_{c.v.} \rho dV + \iint_{c.s.} \rho V \cdot dA$$

# Momentum Equation

- From Newton's 2<sup>nd</sup> Law:
- Net force = time rate of change of momentum

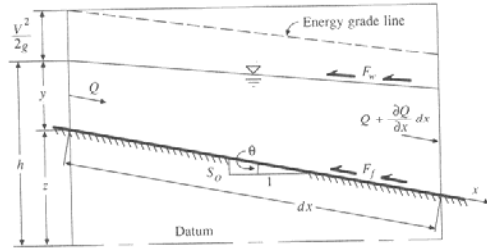
$$\sum F = \frac{d}{dt} \iiint_{c.v.} V \rho dV + \iint_{c.s.} V \rho V \cdot dA$$

Sum of forces on the C.V.

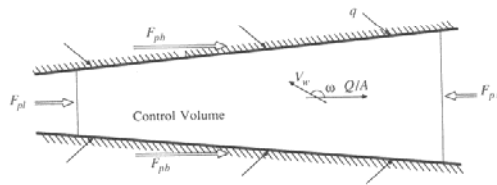
Momentum stored within the C.V.

Momentum flow across the C. S.

### Forces acting on the C.V.



Elevation View

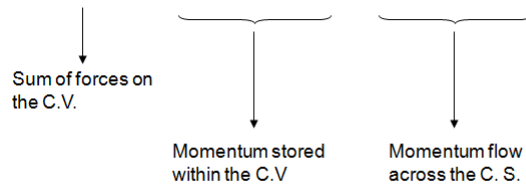


Plan View

- $F_g$  = Gravity force due to weight of water in the C.V.
- $F_f$  = friction force due to shear stress along the bottom and sides of the C.V.
- $F_e$  = contraction/expansion force due to abrupt changes in the channel cross-section
- $F_w$  = wind shear force due to frictional resistance of wind at the water surface
- $F_p$  = unbalanced pressure forces due to hydrostatic forces on the left and right hand side of the C.V. and pressure force exerted by banks

### Momentum Equation

$$\sum F = \frac{d}{dt} \iiint_{c.v.} V \rho dV + \iint_{c.s.} V \rho V \cdot dA$$



$$\frac{1}{A} \frac{\partial Q}{\partial t} + \frac{1}{A} \frac{\partial}{\partial x} \left( \frac{Q^2}{A} \right) + g \frac{\partial y}{\partial x} - g(S_o - S_f) = 0$$



$$\frac{1}{A} \frac{\partial Q}{\partial t} + \frac{1}{A} \frac{\partial}{\partial x} \left( \frac{Q^2}{A} \right) + g \frac{\partial y}{\partial x} - g(S_o - S_f) = 0$$

Local acceleration term	Convective acceleration term	Pressure force term	Gravity force term	Friction force term
-------------------------------	------------------------------------	---------------------------	--------------------------	---------------------------

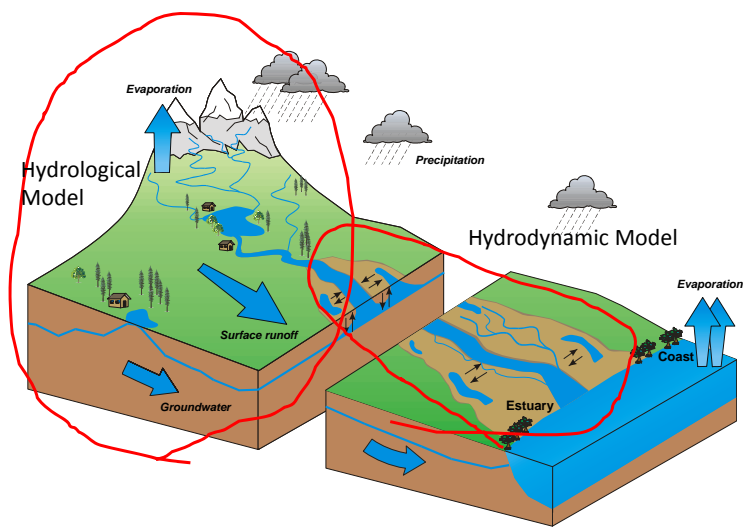
$$\frac{\partial V}{\partial t} + V \frac{\partial V}{\partial x} + g \frac{\partial y}{\partial x} - g(S_o - S_f) = 0$$

Kinematic Wave

Diffusion Wave

Dynamic Wave

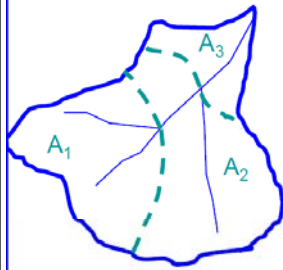
### Flood Modeling Processes that can be simulated



## Hydrological Model Classification

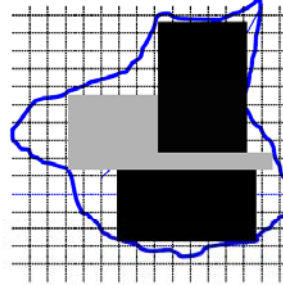
### Lumped

Parameters assigned to each sub-basin



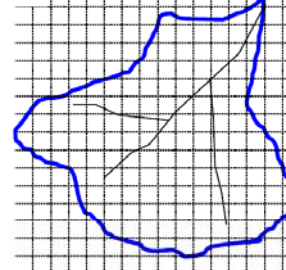
### Semi-Distributed

Parameters assigned to each grid cell, but cells with same parameters are grouped



### Fully-Distributed

Parameters assigned to each grid cell



## BROAD METHODOLOGY

### Stages in the Flood Forecasting

- Computing runoff volume
  - SCS Curve Number Loss
- Modelling direct runoff
  - SCS Transform Method
- Flood Routing
  - Muskingum
- Calibration of the model
- Model validation

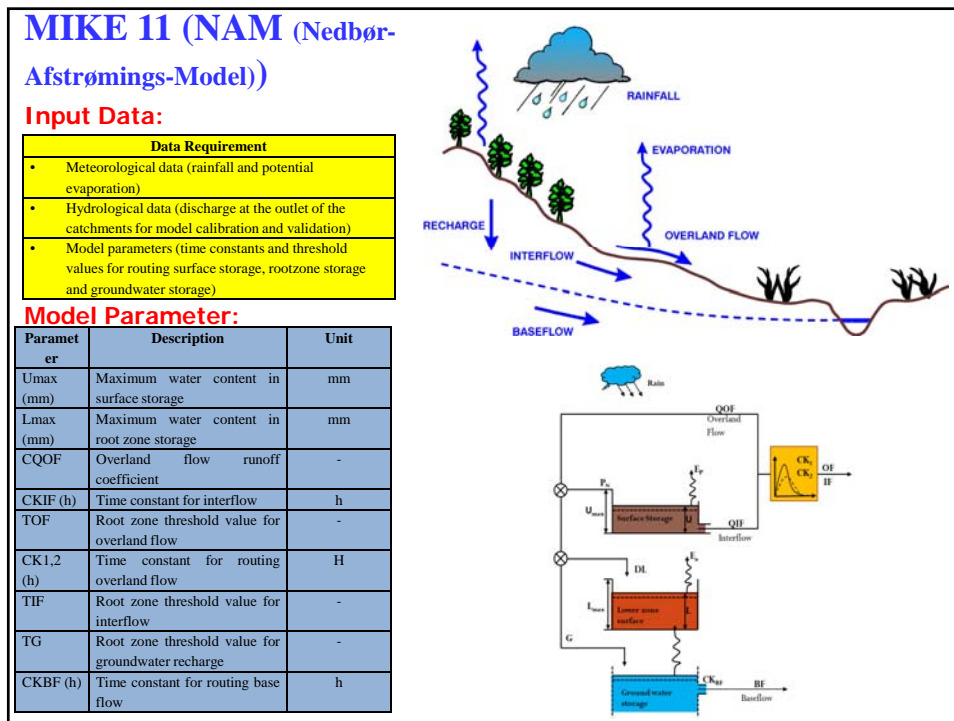
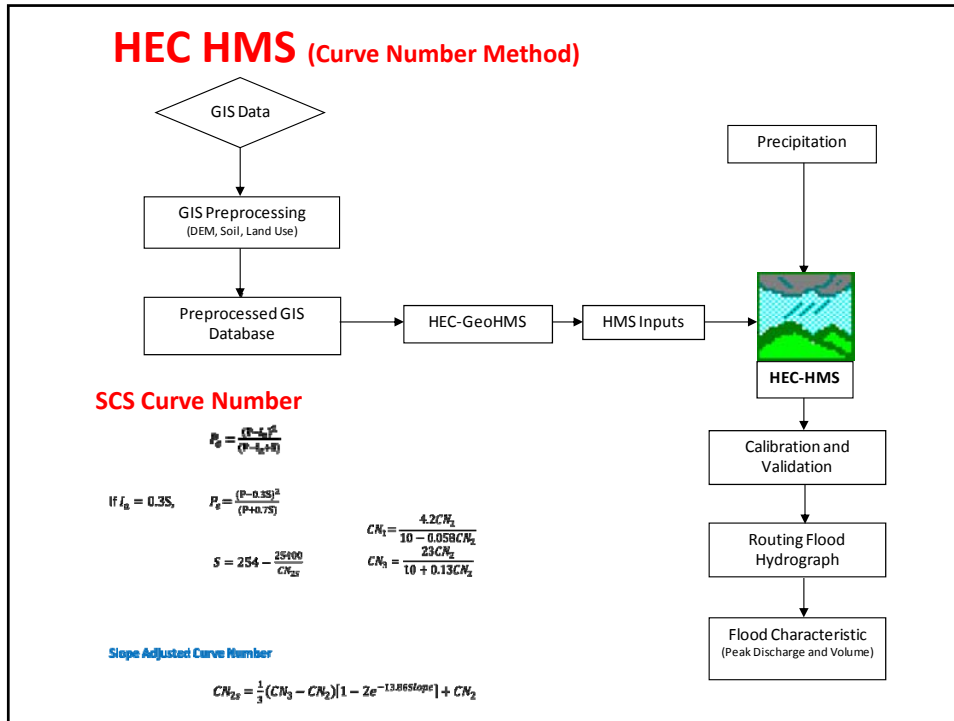
### Spatial and non-spatial Database

#### Static Data Used

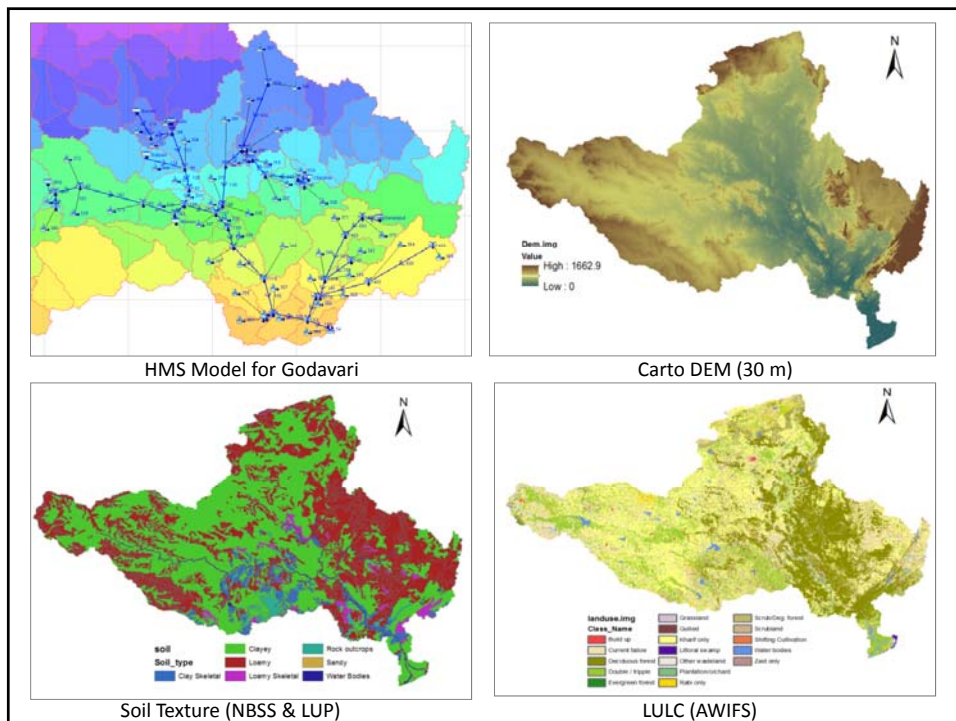
Land use, Soil Texture, CARTO DEM

#### Dynamic Data

- Historic Hydro-meteorological Data
- Real-time 3 hr. hydrological data (CWC)
- Real time In-situ 24 hr Rainfall Data (IMD)
- 24 hr Rainfall Forecast Data at 3 hr frequency (9 km grids from IMD)
- Monthly ET data (computed)

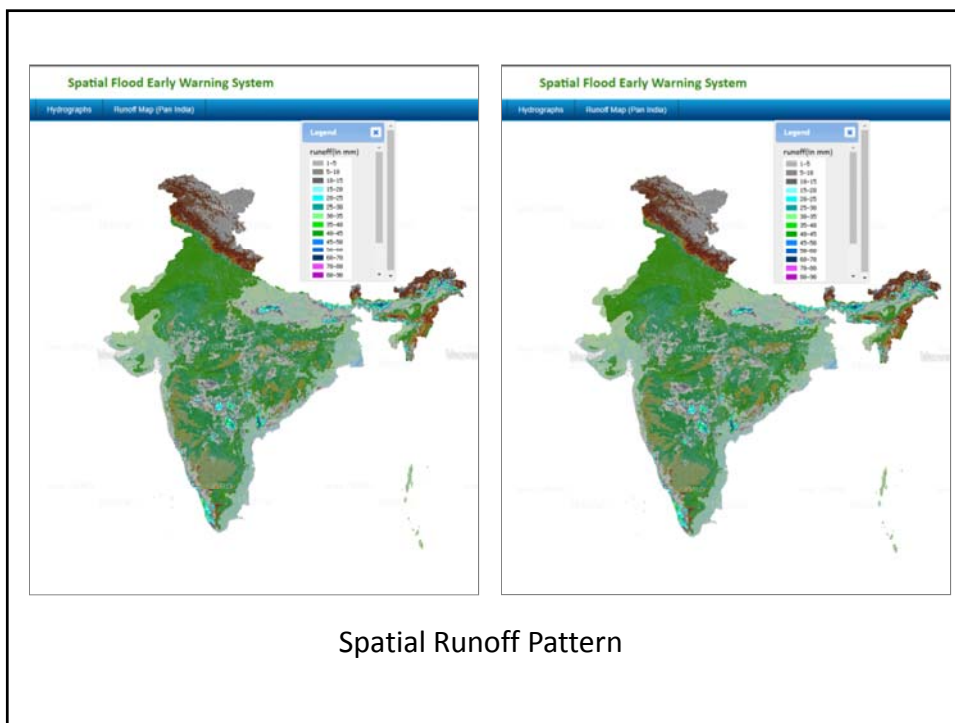
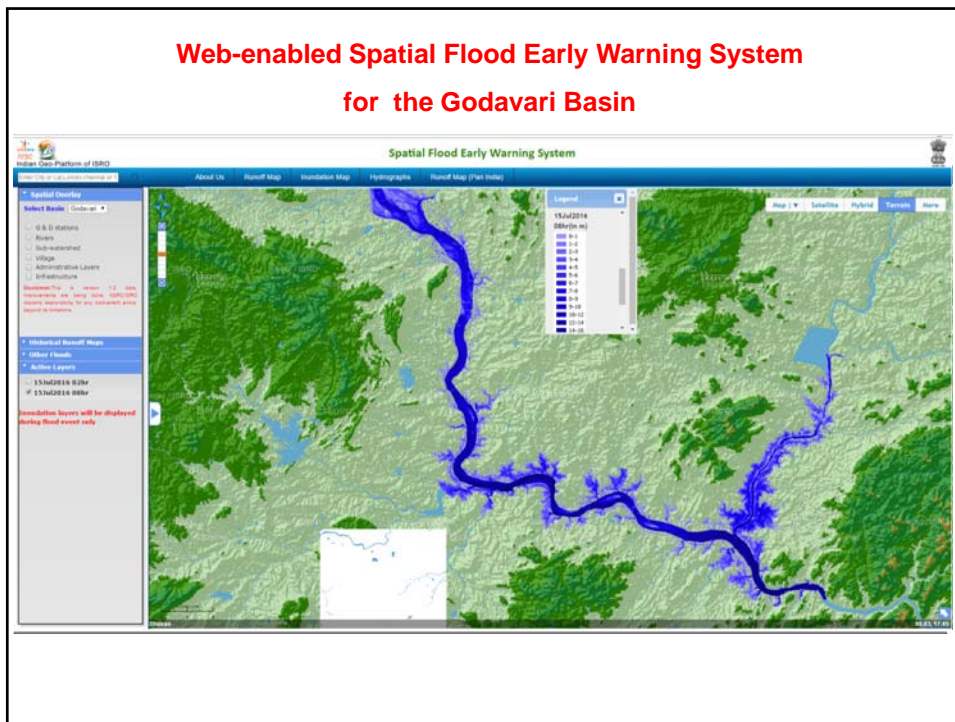


## Case Study of Godavari





### Web-enabled Spatial Flood Early Warning System for the Godavari Basin



Spatial Runoff Pattern

## References

Engineering Hydrology. –Ojha—Oxford University Press 6. Engineering hydrology – K. Subramanyam Tata McGraw Hill, 2009.

Applied Hydrology – Chow, Maidment, Mays, McGraw-Hill

Mays, L.W., Water Resources Engineering, John Willey and Sons, US, 2010.

**Thank You**